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## AN INSTRUMENT FOR THE SPECTROSCOPIC DETERMINATION OF PRECIPITABLE ATMOSPHERIC WATER VAPOR, AND ITS CALIBRATION

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[Weather Bureau, Washington, D. C., January 1940]

The regular daily determination of the amount of precipitable water in the atmosphere is of importance in meteorology for both practical and theoretical purposes. Surface determinations have been made by means of the

The Weather Bureau is now greatly indebted to the Smithsonian Institution which, through the courtesy of Charles G. Abbot and L. B. Aldrich, designed and fabricated for the Weather Bureau an improved type of water-

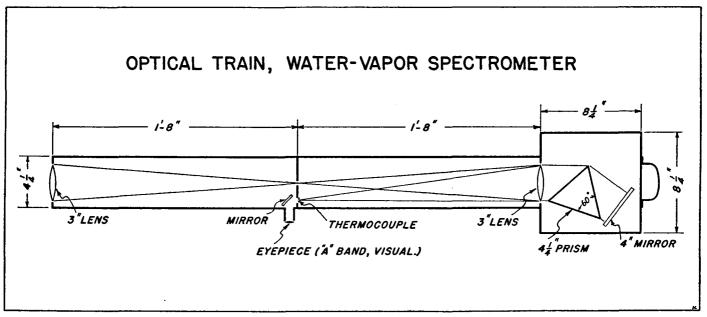


FIGURE 1.—Optical train of the water-vapor spectroscope.

psychrometer ever since the earliest days of meteorology; while the vertical distribution of water vapor has been determined to limited heights by kite, balloon, and airplane flights and, more recently, by means of the radiosonde. Some years ago, the Weather Bureau initiated optical determinations of precipitable water by pyrheliometric measurements of solar radiation through red and yellow filters which cut off long-wave radiation at approximately 0.61  $\mu$  and 0.51  $\mu$ , respectively. However, this method is subject to considerable uncertainty, especially because the filters do not cut off sharply, and because of various assumptions, for example, the adoption of the value 1.3  $\mu$  for the average size of dust particles in the free atmosphere; our own atmospheric dust observations, as well as observations by others, show the impracticability of using a constant factor for all environments.

vapor spectroscope, of design similar to the one used by the Smithsonian Astrophysical Observatory about 8 years ago. We are further indebted to officials of the same Institution in permitting the calibration of the instrument at its two solar constant stations at Burro Mountain, N. Mex., and Table Mountain, Calif. At the first-named station Alfred F. Moore and Alfred D. Froiland more than doubled the usual total number of bolometric runs in order to obtain as many comparisons as possible; they also took turns in keeping the instrument pointed on the sun. Similar cooperation was furnished at Table Mountain by Fred A. Greeley and Stanley Warner.

The water-vapor instrument consists essentially of a spectroscope containing a collimator lens and a 60°-prism in a Littrow mounting, a lens to focus the solar rays on a slit ½ millimeter in width, and a thermocouple having two rectangular surfaces as nearly identical as it is practical to make them. Figure 1 shows the optical train, and figures 3 and 4 an exterior view of the instrument. The thermocouple,

<sup>&</sup>lt;sup>1</sup> Kimball, Herbert H., and Hand, Irving F. The use of glass color screens in the study of atmospheric depletion of solar radiation. Mo. Wes. Rev., March 1933. 61: 80-83. <sup>2</sup> Hand, Irving F. The character and magnitude of the dense dust-cloud which passed over Washington, D. C., May 11, 1934. Mo. Wes. Rev., 62: 186-187. May 1924.

designed and made by Leland B. Clark of the Smithsonian Institution, is of a new and highly efficient design. It is so arranged that by moving it into three different fixed positions, measurements may be made in different parts of the spectral band: The two sections of the solar energy spectrum observed are (A), a  $0.003\mu$  width centered around  $0.9560\mu$ , where the energy curve is not appreciably affected by water-vapor absorption; and (B), a  $0.003\mu$  width in the

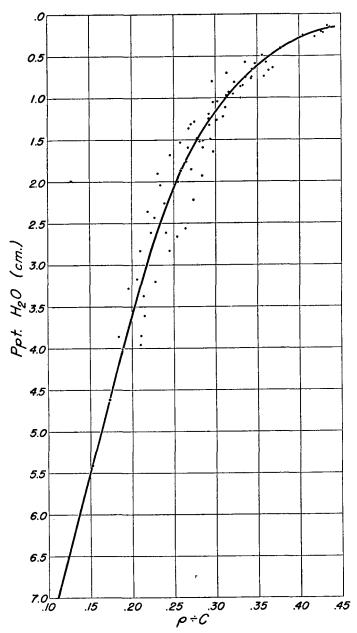


FIGURE 2.—Amounts of precipitable water corresponding to different ratios obtained by readings with the spectroscope.

center of the rho water-vapor absorption band at wavelength  $0.935\mu$ . The base line (C) is obtained by completely shading one strip. A high-sensitivity galvanometer is used for all measurements. In position 1, one strip on A and the other on B, we measure the differential between rho and the crest, or the diminution in energy owing to water-vapor absorption; in position 2, one strip on B and the other completely shaded, we obtain the height of rho above the base line; in position 3, one strip on A and the other completely shaded, we determine the energy at the crest, which should equal the sum of the energies obtained

in the first two positions.<sup>3</sup> This value also may be used for a rough determination of the total normal-incidence radiation.

To eliminate errors arising from inequalities in strip characteristics, the strips are reversed after making the initial reading in each position. Although any two of the positions would suffice for a reading, in actual practice all three positions are read and, as an added check, immediately reread in reverse order to minimize arithmetical errors.

The instrument is kept directed to the sun, manually, by means of a sight at the outer end of the tube through which a pencil of rays passes on to a target at the lower end. A circular opening near the outer end, through which the observer may see when the solar rays are concentrated centrally on the slit, enables him to make an appropriate spot on the target. A small prism projects the portion of the spectral band containing the A line at 0.760 µ to a field which may be viewed by means of an eye-piece containing a cross hair, in order to obtain an approximate setting of the strips on the rho band. The final setting is effected by reading the instrument in position 2 and adjusting the mirror until a minimum reading is obtained. If the instrument is mounted rigidly in a fixed position neither of these adjustments is necessary except at infrequent intervals as checks. Owing to the great length of the tube and the extreme sensitivity of the thermocouple, it is necessary for one observer to continually make both altitude and aximuth adjustments, and at the same time change the position of the thermopile and turn it through 180° upon signal from the observer during readings.

At the two Smithsonian solar-constant stations, complete records of the solar energy curve from  $0.34\mu$  to  $2.34\mu$ are made on each clear day by means of the vacuum bolometer 4; six bolographs are used for the determination of the solar constant on very clear days, and three on less clear days. As previously stated, in order to obtain as many comparisons as possible during these tests, the number of bolographs was increased to as much as three times normal on exceptionally clear days. By reduction of these bolographs the Smithsonian observers have calculated the amount of precipitable water in the atmosphere during each of the runs and these values have been plotted as ordinates against the readings of the new spectroscope as abscissas, fig. 2. Aldrich believes that the Table Mountain results give more accurate values of precipitable water in the atmosphere than the Burro Mountain readings; hence the Table Mountain calibrations have been used exclusively in preparing table 1, which gives the amounts of precipitable water corresponding to different readings of the new instrument.

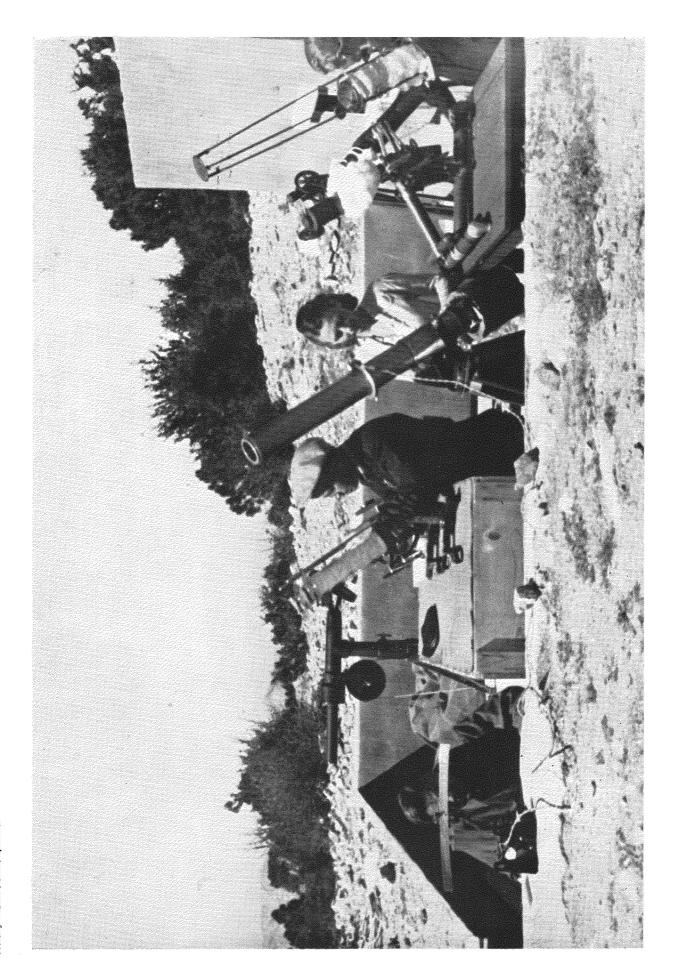
All values obtained at Table Mountain (with the exception of those obtained on the morning of October 20 when the instrument obviously was out of adjustment) have been utilized for the calibration. The probable error of a single observation, calculated from the line of best of

all the readings by the formula  $0.6745\sqrt{\frac{\Sigma v^2}{n-1}}$ , is  $\pm 0.19$  cm.

The individual half-day runs were also plotted, and the probable error of a single observation calculated from the lines of best fit; smooth curves were drawn, rather than lines zig-zagging from point to point. In general the curves of the individual half-day runs are very similar to

<sup>\*</sup> Moore, A. F. Scouting for a site for a solar radiation station. Smithsonian Misc. Coll. Vol. 89, No. 4. Washington, 1933.

4 Abbot, C. G., Fowle, F. E., and Aldrich, L. B. The Vacuum Bolometer. Annals of the Astrophysical Observatory of the Smithsonian Institution, Volume 4, Chapter 2, pp. 45-64. Washington, 1922.



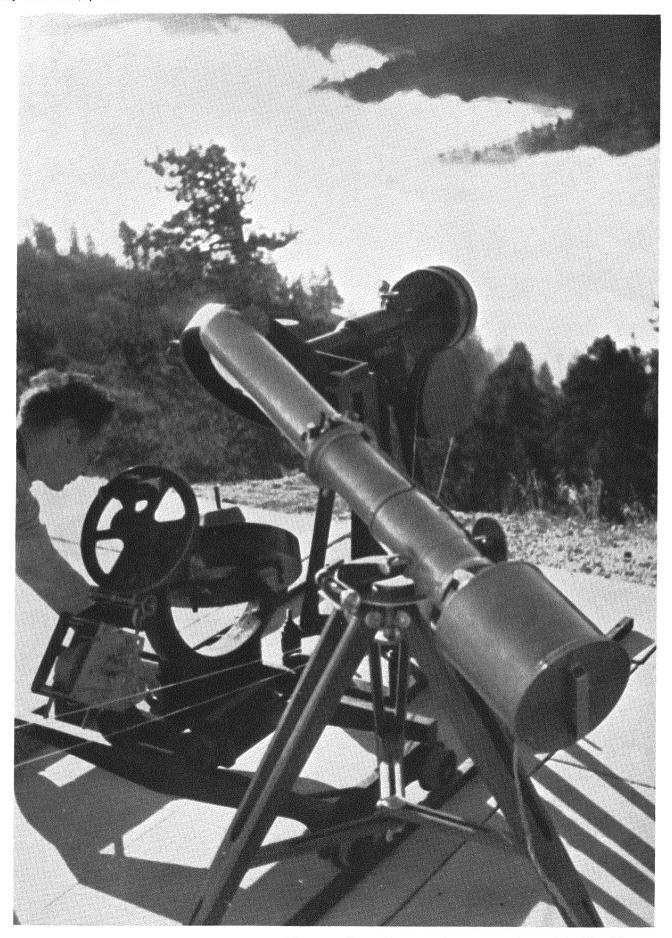


Figure 4.—Water-vapor spectroscope at Table Mountain, Calif. Radiation fog rising up slopes of the Blue Mountains.

the curve of all the observations, falling symmetrically on both sides of the latter; the instrument was adjusted on rho immediately preceding each run, but since it was impossible to exactly duplicate the settings with the temporary set-up of the instrument, each run was made with the strips set on a slightly different portion of the rho band, which explains why the individual days vary even though they closely parallel one another. It has been calculated that one half of the probable error of  $\pm 0.19$  cm. is due directly to these improper adjustments on rho; to eliminate this portion of the error, we shall replace the cross hair so that it bisects the double A line when the best possible setting on rho is made, and the settings may then be effected with great accuracy and with assurance that all are alike. The probable error of future readings should be thus reduced from  $\pm 0.19$  cm. to about  $\pm 0.10$  cm. This reduction is highly important because we shall be concerned chiefly with changes in the amounts of atmospheric water.

Table 2 gives the probable errors of single observations from the individual lines of half-day runs.

Table 1.—Amounts of precipitable water in centimeters corresponding to different ratios of  $\rho/C$  obtained with the water-vapor spectroscope

p/C	em.	ρ/C	cm.	ρ/C	em.	ρ/C	cm.	ρ/C	em.	ρ/C	cm.	ρ/С	em.
0.110 0.111 0.112 0.113 0.114 0.115 0.116 0.117	7. 03 6. 99 6. 96 6. 92 6. 88 6. 85 6. 81 6. 77 6. 74 6. 70	0. 160 0. 161 0. 162 0. 163 0. 164 0. 165 0. 166 0. 167 0. 168 0. 169	5. 13 5. 09 5. 05 5. 01 4. 97 4. 93 4. 89 4. 85 4. 80 4. 76	0. 210 0. 211 0. 212 0. 213 0. 214 0. 215 0. 216 0. 217 0. 218 0. 219	3, 28 3, 25 3, 21 3, 18 3, 15 3, 11 3, 08 3, 04 3, 00 2, 96	0. 260 0. 261 0. 262 0. 263 0. 264 0. 265 0. 266 0. 267 0. 268 0. 269	1. 84 1. 82 1. 80 1. 78 1. 76 1. 74 1. 72 1. 70 1. 68 1. 66	0. 310 0. 311 0. 312 0. 313 0. 314 0. 315 0. 316 0. 317 0. 318 0. 319	1. 02 1. 01 1. 00 . 99 . 98 . 97 . 96 . 95 . 94 . 92	0, 360 0, 361 0, 362 9, 363 0, 364 0, 365 0, 366 0, 367 0, 368 0, 369	0. 54 . 53 . 52 . 52 . 51 . 50 . 50 . 49 . 48 . 48	0. 410 0. 411 0. 412 0. 413 0. 414 0. 415 0. 416 0. 417 0. 418 0. 419	0. 24 . 23 . 23 . 23 . 22 . 22 . 22 . 21 . 21
0.120. 0.121. 0.122. 0.123. 0.124. 0.125. 0.126. 0.127. 0.128.	6. 67 6. 63 6. 59 6. 55 6. 51 6. 18 6. 44 6. 40 6. 37 6. 34	0. 170 9. 171 0. 172 0. 173 0. 174 0. 175 0. 176 0. 177 0. 178 0. 179	4. 72 4. 69 4. 65 4. 62 4. 58 4. 54 4. 51 4. 47 4. 43 4. 40	0. 220 0. 221 0. 222 0. 223 0. 224 0. 225 0. 226 0. 227 0. 228 0. 229	2. 93 2. 90 2. 87 2. 54 2. 81 2. 78 2. 74 2. 71 2. 68 2. 65	0. 270 0. 271 0. 272 0. 273 0. 274 0. 275 0. 276 0. 277 0. 278 0. 279	1. 64 1. 62 1. 60 1. 58 1. 56 1. 55 1. 53 1. 52 1. 50 1. 49	0. 320 0. 321 0. 322 0. 323 0. 324 0. 325 0. 326 0. 327 0. 328 0. 329	. 91 . 90 . 89 . 88 . 87 . 86 . 85 . 84 . 83 . 82	0. 370 0. 371 0. 372 0. 373 0. 374 0. 375 0. 376 0. 377 0. 378 0. 379	. 47 . 46 . 46 . 45 . 44 . 43 . 42 . 42	0. 420 0. 421 0. 422 0. 423 0. 424 0. 425 0. 426 0. 427 0. 428 0. 429	. 20 . 20 . 20 . 19 . 19 . 19 . 19 . 18 . 18
0.130. 0.131. 0.132. 0.133. 0.134. 0.135. 0.136. 0.137. 0.138.	6. 30 6. 26 6. 22 6. 18 6. 14 6. 11 6. 07 6. 04 6. 00 5. 97	0. 180 0. 181 0. 182 0. 183 0. 184 0. 185 0. 186 0. 187 0. 188 0. 189	4. 36 4. 32 4. 28 4. 25 4. 21 4. 18 4. 14 4. 11 4. 07 4. 04	0. 230 0. 231 0. 232 0. 233 0. 234 0. 235 0. 236 0. 237 0. 238 0. 239	2, 62 2, 59 2, 56 2, 53 2, 56 2, 47 2, 45 2, 43 2, 41 2, 38	0. 280 0. 281 0. 282 0. 283 0. 284 0. 285 0. 286 0. 287 0. 288 0. 289	1. 47 1. 45 1. 44 1. 42 1. 40 1. 38 1. 36 1. 34 1. 32 1. 31	0. 330 0. 331 0. 332 0. 333 0. 334 0. 335 0. 336 0. 337 0. 338 0. 339	. 81 . 80 . 79 . 78 . 77 . 76 . 75 . 74 . 73	0, 380 0, 381 0, 382 0, 383 0, 384 0, 385 0, 386 0, 387 0, 388 0, 389	. 41 . 40 . 40 . 39 . 38 . 38 . 37 . 36 . 36 . 35	0. 430 0. 431 0. 432 0. 433 0. 434 0. 435 0. 436 0. 437 0. 438 0. 439	. 18 . 18 . 17 . 17 . 17 . 17 . 16 . 16
0.140_ 0.141_ 0.142_ 0.143_ 0.144_ 0.145_ 0.146_ 0.147_ 0.148_ 0.149_	5. 93 5. 89 5. 85 5. 81 5. 77 5. 73 5. 69 5. 65 5. 61 5. 57	0. 190 0. 191 0. 192 0. 193 0. 194 0. 195 0. 196 0. 197 0. 198 0. 199	4.00 3.96 3.93 3.90 3.87 3.85 3.76 3.76 3.73	0. 240 0. 241 0. 242 0. 243 0. 244 0. 245 0. 246 0. 247 0. 248 0. 249	2. 36 2. 33 2. 30 2. 27 2. 25 2. 22 2. 20 2. 17 2. 14 2. 12	0. 290 0. 291 0. 292 0. 293 0. 294 0. 295 0. 296 0. 297 0. 298 0. 299	1. 29 1. 28 1. 27 1. 26 1. 25 1. 24 1. 22 1. 20 1. 19 1. 17	0. 340 0. 341 0. 342 0. 343 0. 345 0. 345 0. 346 0. 347 0. 348 0. 349	.71 .70 .69 .68 .67 .66 .65 .64	0. 390 0. 391 0. 392 0. 393 0. 394 0. 395 0. 396 0. 397 0. 398 0. 399	. 34 . 33 . 33 . 32 . 32 . 31 . 30 . 30 . 29	0. 446 0. 447 0. 448	. 16 . 16 . 16 . 15 . 15 . 15 . 15 . 15
0.150 0.151 0.152 0.153 0.154 0.155 0.156 0.157 0.158 0.159	5. 54 5. 50 5. 26 5. 42 5. 38 5. 34 5. 30 5. 26 5. 22 5. 18	0. 200 0. 201 0. 202 0. 203 0. 204 0. 205 0. 206 0. 207 0. 208 0. 209	3. 65 3. 62 3. 58 3. 55 3. 51 3. 48 3. 44 3. 40 3. 36 3. 32	0. 250 0. 251 0. 252 0. 253 0. 254 0. 255 0. 256 0. 257 0. 258 0. 259	2. 09 2. 07 2. 04 2. 02 2. 00 1. 97 1. 94 1. 92 1. 89 1. 87	0. 300 0. 301 0. 302 0. 303 0. 304 0. 305 0. 306 0. 307 0. 308 0. 309	1. 15 1. 14 1. 12 1. 11 1. 10 1. 08 1. 07 1. 06 1. 04 1. 03	0. 350 0. 351 0. 352 0. 353 0. 354 0. 355 0. 356 0. 357 0. 358 0. 359	. 63 . 63 . 61 . 60 . 60 . 59 . 58 . 57 . 56 . 55	0. 400 0. 401 0. 402 0. 403 0. 404 0. 405 0. 406 0. 407 0. 408 0. 409	. 28 . 28 . 27 . 26 . 25 . 25 . 25 . 25 . 24 . 24	0.450	. 15

Table 2.—Probable errors of single observations from individual curves

Date	Probable error ±	Date	Probable error ±
1939	cm.	1939	cm.
Oct. 17, a. m.	0.14	Oct. 20, p. m	0.11
Oct. 18, a. m	. 10	Oct. 21, a. m.	. 07
Oct. 19, a. m	, 10	Oct. 21, p. m	. 07
Oct. 19, p. m	.05	Oct. 22, p. m	. 07
Oct. 20, a. m.	. 04	Mean	1.08

<sup>&</sup>lt;sup>1</sup>Probable error from plot of all observations, as shown in text, is  $\pm 0.19$  cm.

According to Harrison 5 the average amounts of precipitable water in a vertical atmospheric column over Washington, D. C., are as follows: Spring, 1.69 centimeters; summer, 3.49 centimeters; autumn, 2.23 centimeters; and winter, 1.00 centimeters. As all of our observations will be made through air masses greater than 1.0, the actual amounts of precipitable water measured will in general exceed these values. We found more consistent readings with larger air masses than with the sun near the zenith, which is explained in part by the fact that incipient clouds passing between the sun and instrument create much larger variations during a series when the sun is nearly overhead than when the sun is at a lower altitude, because the incipient cloud with its water content represents a larger percentage of the total water being measured and its slightest movement, unless it is absolutely uniform, alters the readings.

In table 3 are tabulated the original observations together with the times, air masses, reductions, factors, amounts of precipitable water through the observed air masses and the precipitable water reduced to unit air mass.

Table 3.—Reduction of water-vapor observations

BURRO MOUNTAIN, N. MEX.

SEPTEMBER 23, 1939

Time	Air mass	Readings			Factor	Ppt. H <sub>2</sub> O	Ppt. H <sub>2</sub> O 1.0 air mass
8:00 a. m. 8:14 a. m. 8:30 a. m. 8:49 a. m. 9:14 a. m. 9:33 a. m. 10:12 a. m. 11:10 a. m.	2. 47 2. 22 2. 02 1. 81 1. 61 1. 51 1. 35 1. 22	209 204 204 201 200 203 199 193	101 106 109 109 114 123 126 128	310 309 314 310 318 325 327 327	0. 326 . 341 . 349 . 352 . 367 . 376 . 390 . 404	Cm. 1. 35 1. 19 1. 10 . 99 . 89 . 90 . 79 . 63	Cm 0. 55 . 54 . 54 . 55 . 55 . 60 . 59 . 52
SEP	ТЕМВІ	ER 24,	1939				
6:48 a. m. 6:58 a. m. 7:10 a. m. 7:29 a. m. 7:52 a. m. 9:37 a. m. 9:36 a. m. 11:04 a. m.	6. 18 5. 14 4. 21 3. 31 2. 52 1. 51 1. 40 1. 23	230 231 227 226 227 205 205 199	50 59 65 75 84 101 102 107	279 290 299 301 315 304 313 315	0. 177 . 203 . 232 . 249 . 275 . 328 . 339 . 359	5. 12 4. 50 3. 81 2. 85 2. 23 1. 51 1. 49 1. 25	0.83 .88 .90 .86 .88 1.00 1.06
SEP.	гемве	R 25,	1939				<del></del>
6:46 a. m. 6:58 a. m. 7:07 a. m. 7:18 a. m. 7:24 a. m. 8:01 a. m. 8:17 a. m. 8:39 a. m.	6. 96 5. 39 4. 51 3. 96 3. 53 2. 51 2. 20 1. 93 1. 51	276 279 278 274 273 262 256 250 235	40 48 56 67 65 81 89 93 110	317 329 332 342 341 347 347 351 353	0. 128 . 149 . 165 . 198 . 196 . 241 . 260 . 280 . 327	Off tal 5. 93 4. 50 3. 81 2. 85 2. 23 1. 51 1. 49 1. 25	ple. 1. 10 1. 00 . 95 . 81 . 89 . 69 . 77 . 83
SEP	темві	FR 26,	1939				
6:40 a. m. 6:48 a. m. 7:00 a. m. 7:12 a. m. 7:30 a. m. 8:00 a. m. 8:32 a. m. 9:35 a. m.	7. 87 6. 51 5. 14 4. 23 3. 32 2. 50 2. 00 1. 50 1. 27	261 260 259 256 249 243 230 224 204	35 41 50 55 63 75 84 100 110	297 301 309 313 319 320 316 326 316	0. 120 . 136 . 162 . 179 . 211 . 238 . 270 . 311 . 352	Off te 6. 55 5. 40 4. 72 3. 67 2. 87 2. 27 1. 68 1. 39	1.01 1.05 1.12 1.11 1.15 1.14 1.11 1.09

<sup>&</sup>lt;sup>1</sup> Harrison, Louis P. On the water-vapor in the atmosphere over the United States east of the Rocky Mountains. Mo. Wea. Rev. 59: 470. December, 1931.

Table 3.—Reduction of water-vapor observations—Continued SEPTEMBER 27, 1939

5211	EMIDE	R 27, 1939			
Time	Air mass	Readings	Factor	Ppt. H <sub>2</sub> O	Ppt. H <sub>2</sub> O 1.0 air mass
34 a. m	9. 30 6. 76 5. 59 4. 77 4. 02 2. 52 2. 00 1. 50 1, 23	259 36 294 266 44 312 259 50 311 258 57 316 256 66 323 238 87 329 222 112 343 217 114 337 202 130 339	0. 121 . 145 . 165 . 182 . 206 . 272 . 344 . 350 . 398	Cm. Off tab 6. 18 5. 35 4. 71 3. 81 2. 13 1. 27 1. 05 . 79	Cm. e. 0.91 .96 .98 .95 .85 .64 .70
:39 p. m :45 p. m :07 p. m :27 p. m :43 p. m :57 p. m :57 p. m :58 p. m :518 p. m :528 p. m	4. 25	227 145 372 232 109 349 237 99 344 238 97 344 245 90 345 243 84 341 239 75 321 244 64 314 246 57 309	.390 .313 .303 .299 .279 .269 .247 .215 .196	. 76 1. 21 1. 52 1. 66 1. 89 2. 27 2. 49 3. 36 4. 24	. 57 . 57 . 61 . 56 . 54 . 53 . 48 . 53
SEP	темве	ER 28, 1939			
6:39 a. m 6:40 a. m 7:03 a. m 7:14 a. m 7:35 a. m 8:04 a. m 8:36 a. m 9:41 a. m 11:25 a. m	6. 72 5. 05 4. 18 3. 28 2. 51 2. 01 1. 50	234 85 323 233 92 325 233 104 333 226 109 337 217 118 339 212 123 340 205 128 340 200 136 332 190 137 335	. 283 . 304 . 327 . 356 . 372 . 391 . 401	2. 23 1. 89 1. 34 1. 05 . 85 . 69 . 59 . 48 . 42	. 26 . 28 . 27 . 25 . 26 . 27 . 29 . 32 . 34
1:33 p. m. 3:17 p. m. 3:32 p. m. 3:34 p. m. 3:46 p. m. 4:07 p. m. 4:36 p. m. 4:35 p. m. 5:18 p. m. 5:18 p. m.	1. 84 2. 00 2. 16 2. 50 3. 26 4. 15 5. 02 6. 24	195 128 331 208 123 337 214 106 331 213 109 329 220 103 327 231 89 324 236 79 318 240 69 306 236 66 306 231 56 302	.323 .283 .254 .223 .224	. 56 . 55 . 98 . 84 1. 19 1. 70 2. 53 3. 26 3. 51 4. 45	. 42 . 30 . 48 . 38 . 48 . 55 . 61 . 65 . 56
<sup>1</sup> Clouds during observation. SEP	темві	ER 29, 1939	•	•	··
6:37 a. m	8.87 6.96 5.12 4.21 3.30 2.51 2.01	234 78 322 234 83 324 234 95 333 222 94 321 220 95 322 213 103 324 203 113 325 192 125 323	. 272 . 293 . 303 . 310 . 335 . 370	2. 52 2. 02 1. 59 1. 40 1. 20 . 97 . 73 . 58	0. 28 29 31 .33 .36 .39 .36 .39
0.0	70000	TAIN, CALIF.	_ [ _	<u> </u>	<u>.l</u>
6:56 a. m. 7:12 a. m. 7:37 a. m. 8:03 a. m. 8:41 a. m. 10:05 a. m. 11:42 a. m.	4. 90 3. 88 2. 99 2. 50 2. 01 1. 50	205 81 300 186 77 27 184 89 276 179 100 27 175 98 27 173 90 26 182 79 256	309 330 358 358 364 347	1. 49 1. 22 . 86 . 74 . 67 . 59 . 80	0. 30 . 31 . 20 . 30 . 33 . 38 . 57
	TOBEI	3. 18, 1939	1	<del></del>	1
1.43 p. m. 2:05 p. m. 2:08 p. m. 2:38 p. m. 2:49 p. m. 3:08 p. m. 3:08 p. m. 3:08 p. m. 3:58 p. m. 4:14 p. m. 4:24 p. m. 4:34 p. m. 4:34 p. m.	1.80 1.99 2.20 2.49 3.00 4.00 4.99 5.93	196 81 277 192 71 267 211 79 283 210 75 277 208 68 270 215 64 277 214 60 274 211 56 267 212 53 260 213 48 263	7 .275 3 .268 9 .258 6 .246 9 .231 4 .219 7 .210	1. 28 1. 36 1. 53 1. 68 1. 90 2. 36 2. 83 3. 17	0. 66 . 76 . 68 . 71 . 67 . 67 . 55 . 55

Table 3.—Reduction of water-vapor observations—Continued OCTOBER 19, 1939

1	OBER					Ppt.
Time	Air mass	Rea	dings	Factor	Ppt. H₃O	H <sub>2</sub> O 1.0 air mass
3:36 a. m. 3:56 a. m. 7:12 a. m. 7:39 a. m. 8:02 a. m. 8:40 a. m. 10:07 a. m.	7. 41 4. 97 4. 00 3. 00 2. 50 2. 01 1. 50 1. 39	245 238 237 217 224 224 202 212	75 320 96 338 98 339 99 322 101 328 109 327 112 318 110 324	.314 .336 .361	Cm. 2.04 1.33 1.05 .83 .70 .57 .57	Cm. 0. 28 . 27 . 26 . 28 . 28 . 28 . 38 . 47
12:57 p. m 2:26 p. m 3:04 p. m 3:33 p. m	1. 50 2. 00 2. 49 3. 09 4. 02	186 193 192 193 197	94 286 88 283 83 278 74 279 68 271	.316 .302 .293	. 77 . 93 1. 04 1. 32 1. 76	.51 .46 .42 .43
00	TOBER	20, 19	39			
1·27 p. m. 2·27 p. m. 3·04 p. m. 3·27 p. m. 3·54 p. m. 4·09 p. m. 4·19 p. m.	1. 61 2. 02 2. 41 3. 00 4. 03 5. 01 5. 99 7. 44	207 209 210 213 214 212 211 215	108 311 99 300 89 300 83 300 75 300 67 289 64 279 59 270	321 302 302 285 274 254 2 228	0.75 .95 1.26 1.59 2.22 2.66 3.20 3.96	0. 47 . 47 . 52 . 53 . 55 . 53 . 53 . 53
ос	TOBER	21, 19	39			
6:32 a. m. 6:42 a. m. 6:58 a. m. 7:14 a. m. 7:42 a. m. 8:05 a. m. 8:45 a. m. 11:34 a. m.	9. 35 7. 21 4. 98 3. 98 3. 02 2. 50 2. 01 1. 41	213 213 207 206 218 202 193 184	55 27. 67 28 78 29 84 29 99 31 99 30 111 30 109 29	1 .245 3 .284 6 .297 7 .312 5 .333 8 .369	3.85 2.83 1.93 1.64 1.11 .85 .64	0. 41 . 39 . 39 . 41 . 37 . 34 . 32 . 30
1:08 p. m	2, 98 3, 29 3, 99	188 190 202 206 206 207 211 219	100 29 93 28 82 28 76 28 71 28 69 28 60 27 50 27	9 .336 7 .292 3 .271 5 .267 0 .255 5 .227	0. 49 .75 1. 24 1. 31 1. 59 2. 00 2. 43 3. 28	0. 32 . 38 . 42 . 40 . 40 . 41 . 46
00	TOBER	22, 19	939			
6:34 a. m. 6:44 a. m. 7:00 a. m. 7:16 a. m. 7:44 a. m. 8:07 a. m.	5.02	215 214 209 208 197 204	57 27 67 28 75 28 80 29 93 29 90 30	3 .241 9 .271 4 .285 8 .330	3. 61 2. 61 1. 85 1. 51 1. 00 . 97	0. 41 . 38 . 37 . 37 . 33 . 39
00	TOBER	23, 19	939			
1:28 p. m. 1:42 p. m. 2:29 p. m. 2:59 p. m. 3:21 p. m. 3:50 p. m. 4:05 p. m. 4:16 p. m. 4:26 p. m. 4:37 p. m.	2. 50 2. 99 3. 97 4. 98 5. 93 7. 30	208 208 206 216 223 221 224 224 222 185	76 29 77 29 66 28 63 28 61 29 58 26 50 27 47 27 39 26 23 20	4 .281 4 .259 9 .239 0 .223 3 .214 8 .188 1 .173 3 .153	1. 48 1. 52 1. 87 2. 26 2. 61 3. 37 4. 01 4. 61 5. 40 7. 00	0. 89 . 88 . 89 . 90 . 87 . 85 . 81 . 78 . 74
00	товен	25, 1	939			
7:47 a. m	2. 51 2. 21 2. 01 1. 80	184 131 175 176 177 181	127 32 178 31 134 31 131 31 130 31 120 30	5 .429 3 .437 5 .434 1 .247	0. 27 . 22 . 16 . 14 . 21 . 25	0. 09 . 09 . 07 . 07 . 12 . 14
Mean	<u> </u>				-	. 466

<sup>\*</sup>The factors are obtained from the following formula:  $\frac{1}{2}(AC+AB-AC)/\frac{1}{2}(AB\frac{1}{2}BC+AC)$  where A=C rest; B the rho band and C the base line.